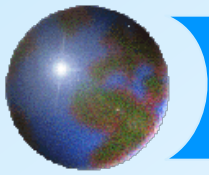


Global Energy Balance

GEOG/ENST 2331: Lecture 4

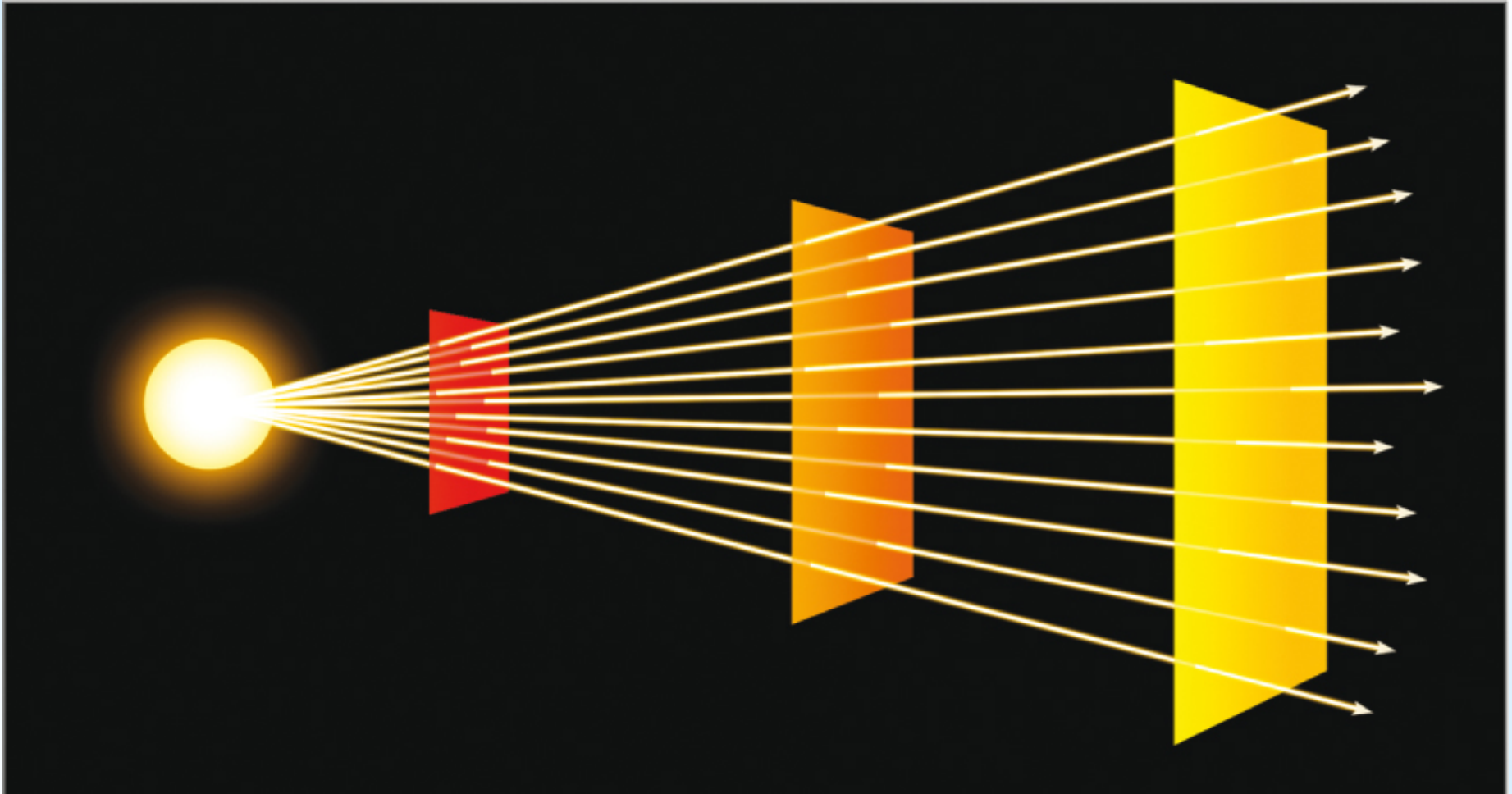
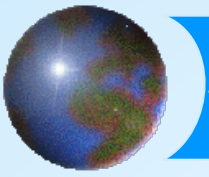
Ahrens: Chapter 2

Lab 1

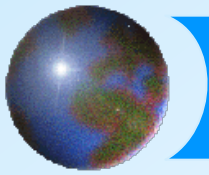


Absorption and Transmission

- ⊕ When radiation reaches the atmosphere, it can be scattered, reflected, absorbed or transmitted
 - ⊞ *Albedo* determines how much is reflected/scattered
 - ⊞ *Absorptivity* determines how much of what is left is absorbed or transmitted
- ⊕ *Black bodies* absorb all non-reflected radiation
- ⊕ *Selective absorbers* absorb only specific wavelengths, remainder is transmitted

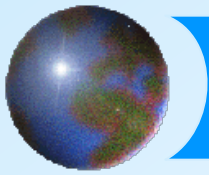


Ahrens: Ch. 2 Fig. 2



Total Solar Irradiance (S)

- ✦ *Quantity* of electromagnetic radiation is not reduced with distance through a vacuum
- ✦ *Intensity* is reduced as energy becomes distributed over a larger area
- ✦ Therefore, radiation intensity decreases in proportion to the square of the distance



Irradiance

☉ *Inverse square law*

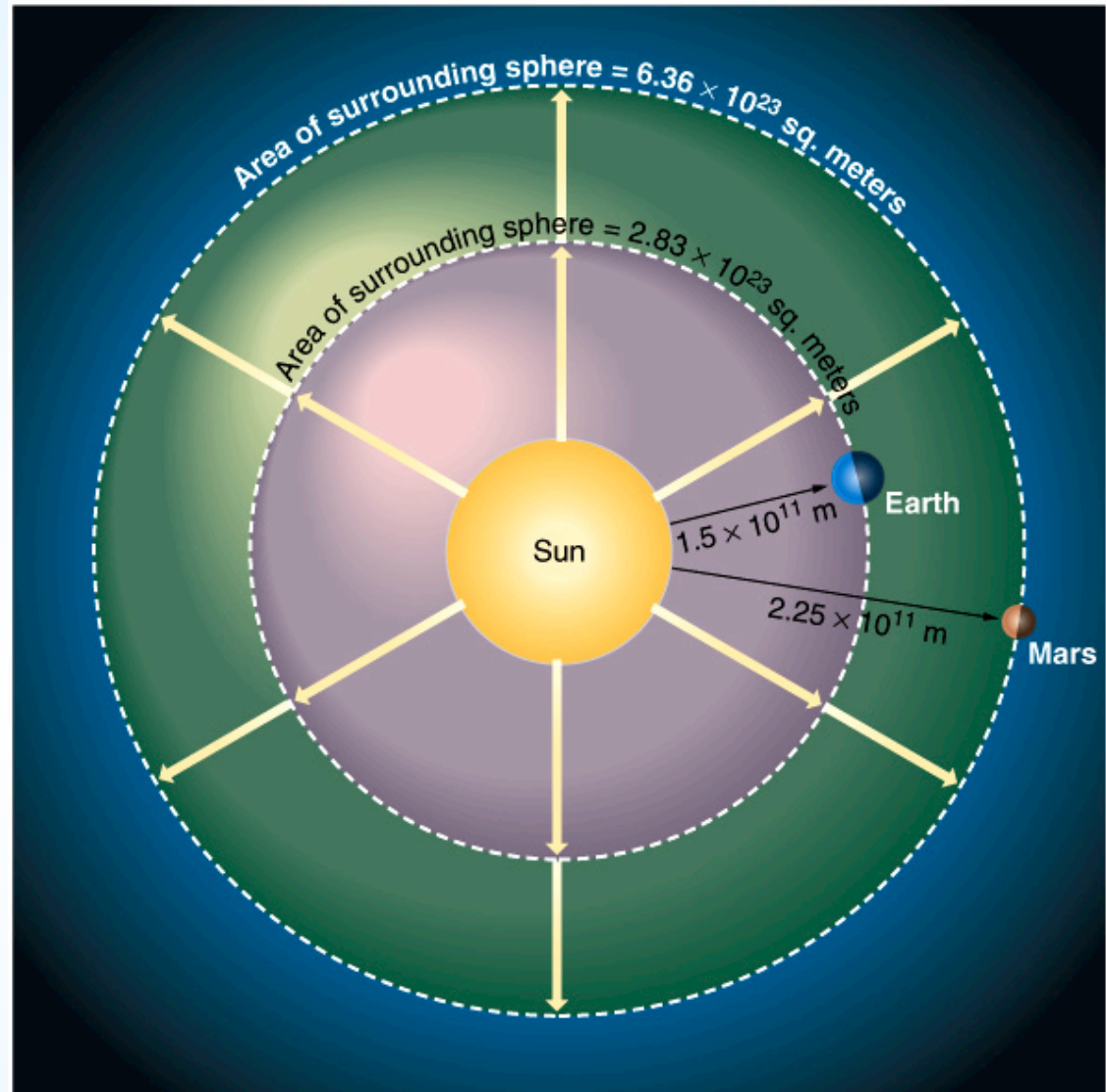
S is proportional to $\frac{1}{d^2}$

☉ **Earth:**

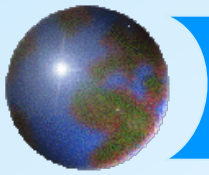
$$S = 1380 \text{ Wm}^{-2}$$

☉ **Mars:**

$$S = 445 \text{ Wm}^{-2}$$



A&B: Figure 2.9



Area of interception

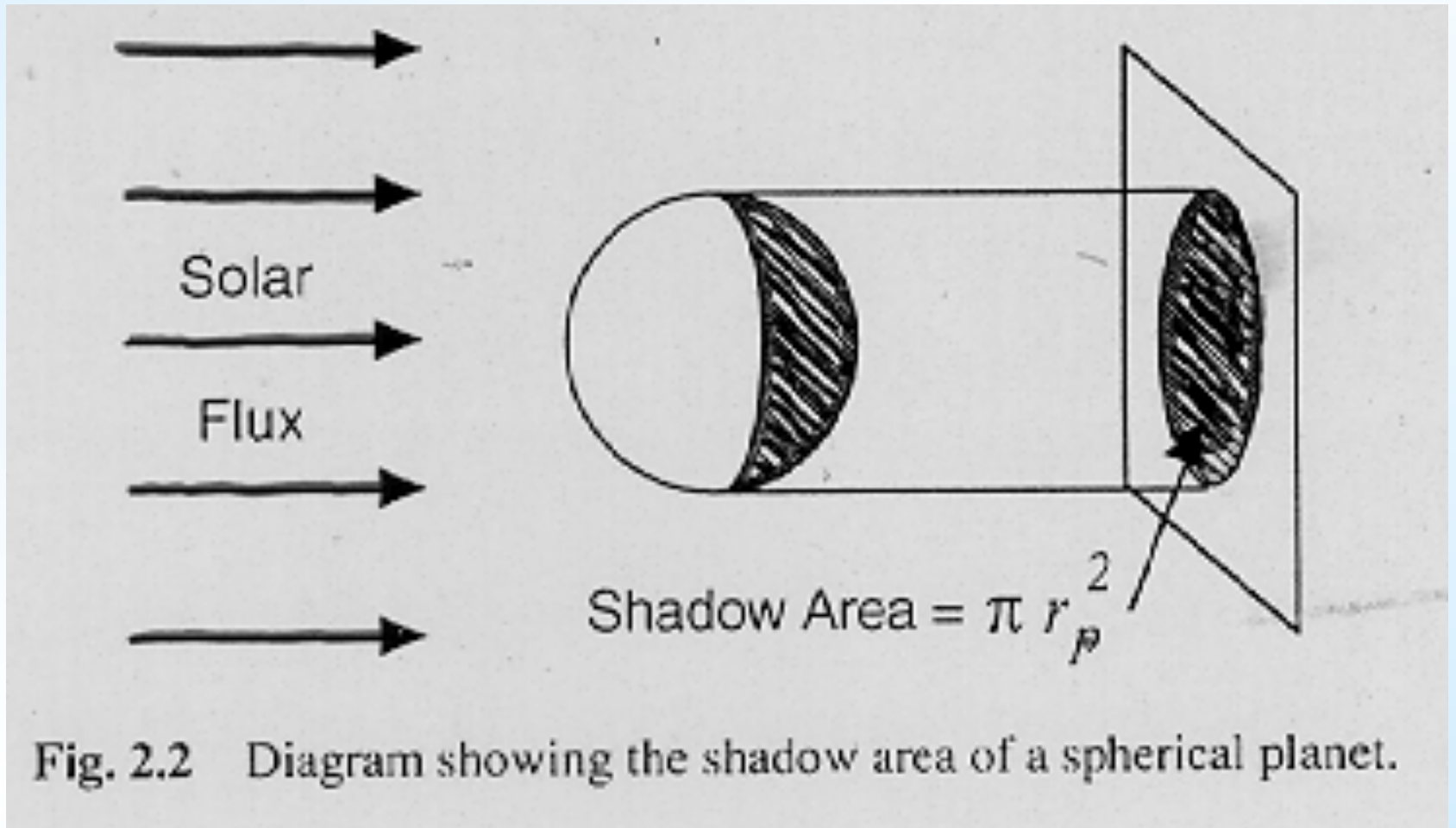
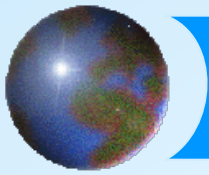
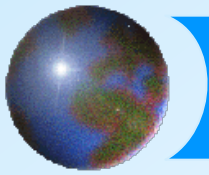


Fig. 2.2 Diagram showing the shadow area of a spherical planet.



First Law of Thermodynamics

- ⊕ Energy cannot be created or destroyed
- ⊕ If a system is at equilibrium, the amount of energy coming in must be equal to the amount of energy going out.



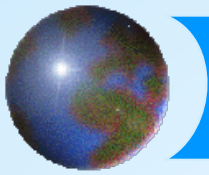
Total solar irradiance

✚ 1380 W m⁻²

- ✚ Incoming shortwave radiation
- ✚ 30% reflected, 70% is absorbed

✚ $1380 \times 0.7 = 967 \text{ Wm}^{-2}$

- ✚ *Absorbed* incoming radiation
- ✚ Must be matched by outgoing IR radiation



Incoming and Outgoing

☉ Incoming radiation

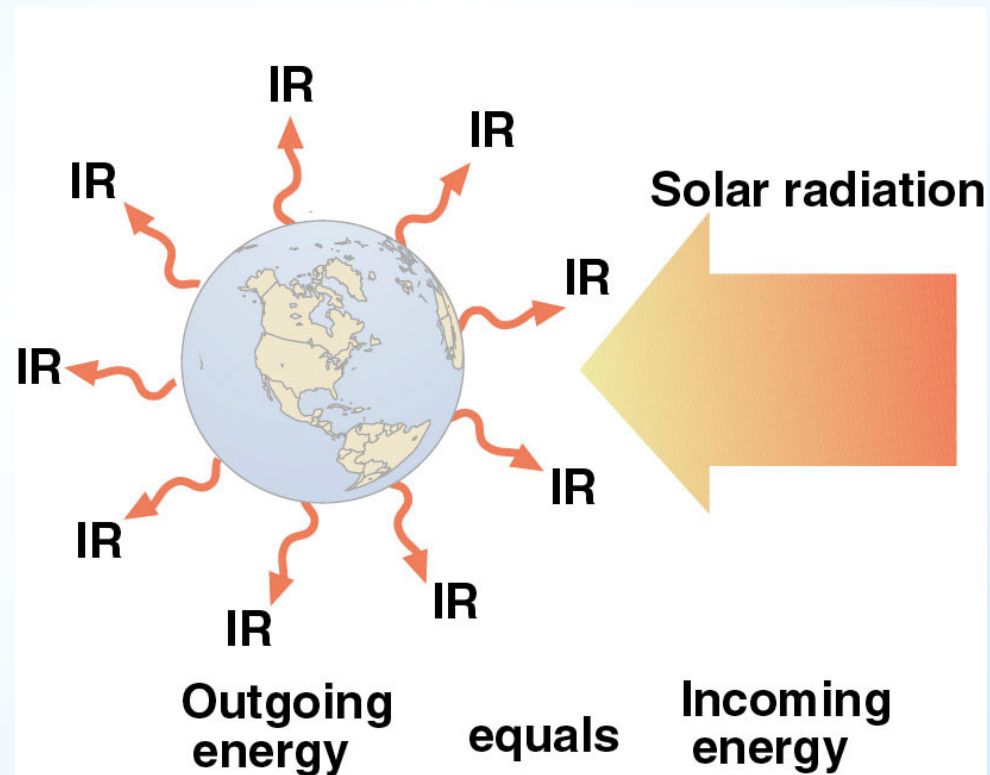
- ☒ Intercepted on a circle

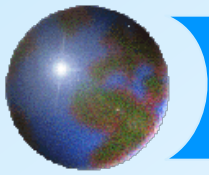
$$\text{area} = \pi r_e^2$$

☉ Outgoing radiation

- ☒ Radiated over a sphere

$$\text{area} = 4\pi r_e^2$$





Energy Balance

$$\text{Incoming} = 967 \times \pi r_e^2$$

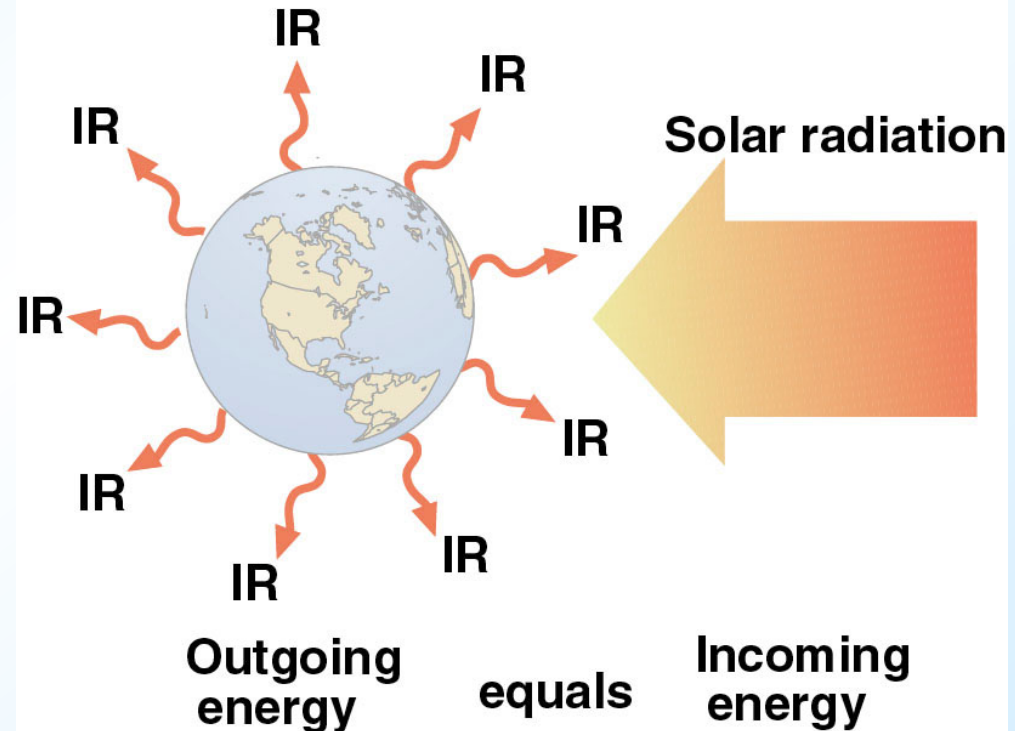
$$\text{Outgoing} = \sigma T_e^4 \times 4\pi r_e^2$$

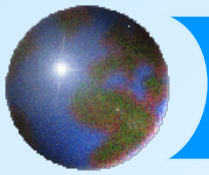
(Earth is a blackbody)

$$\text{Therefore: } \sigma T_e^4 = 967/4$$
$$= 242 \text{ Wm}^{-2}$$

Solve to get $T_e = 255 \text{ K}$

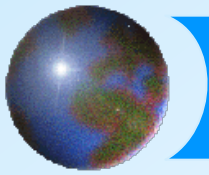
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Radiative Equilibrium Temperature

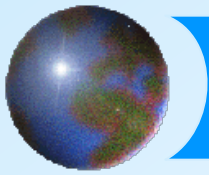
- ⊕ Longwave emission matches shortwave absorption
- ⊕ If you measured the Earth's outgoing radiation from space, this would be the temperature you would calculate
- ⊕ $255 \text{ K} = -18^\circ\text{C}$



Exchange with the atmosphere

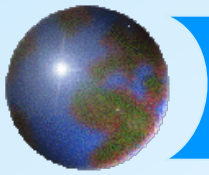
- ✚ Objects radiate in all directions
 - ✚ Earth and Sun radiate outward from all around their spherical surface
 - ✚ Atmosphere is a hollow sphere; radiates both out (up) and in (down)

- ✚ There is an exchange of radiation between Earth and atmosphere



Greenhouse calculation

- ✚ Model the atmosphere as one thin layer that:
 - ✚ Absorbs 10% of the incoming solar radiation
 - ✚ Absorbs 80% of the outgoing terrestrial radiation
- ✚ For the entire sphere, let:
 - ✚ x be the radiation emitted from the surface,
 - ✚ y be the radiation emitted from the atmosphere,
and
 - ✚ I be non-reflected radiation entering the top of the atmosphere from the Sun.



I: non-reflected incoming radiation

Total solar irradiance

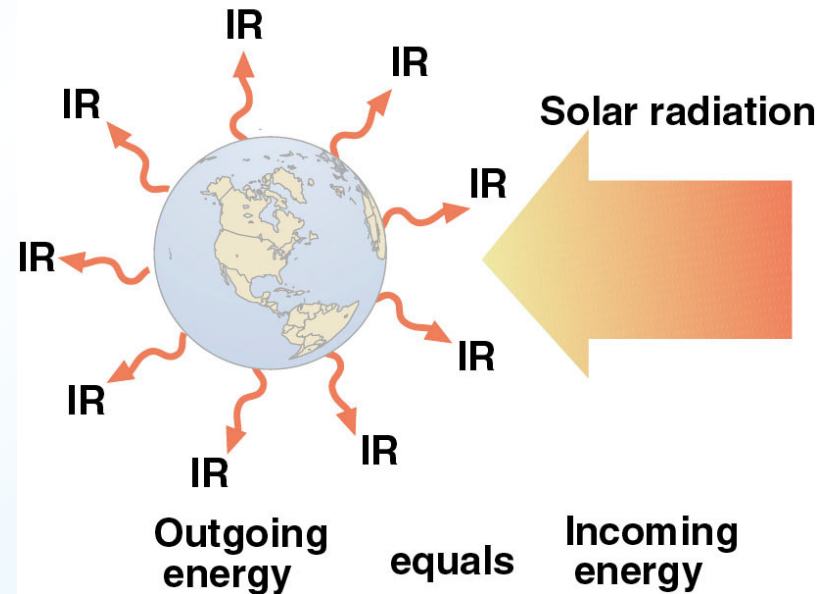
$$S = 1380 \text{ W m}^{-2}$$

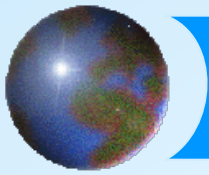
The average albedo $A = 0.3$

Ratio of area of absorption (circle) to area of emission (sphere) is $1/4$

$$I = S(1-A)/4 = 967/4 = 242 \text{ Wm}^{-2}$$

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Greenhouse calculation

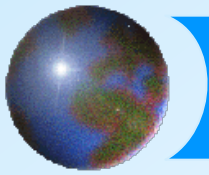
What do we know about x and y ?

Stefan-Boltzmann law:

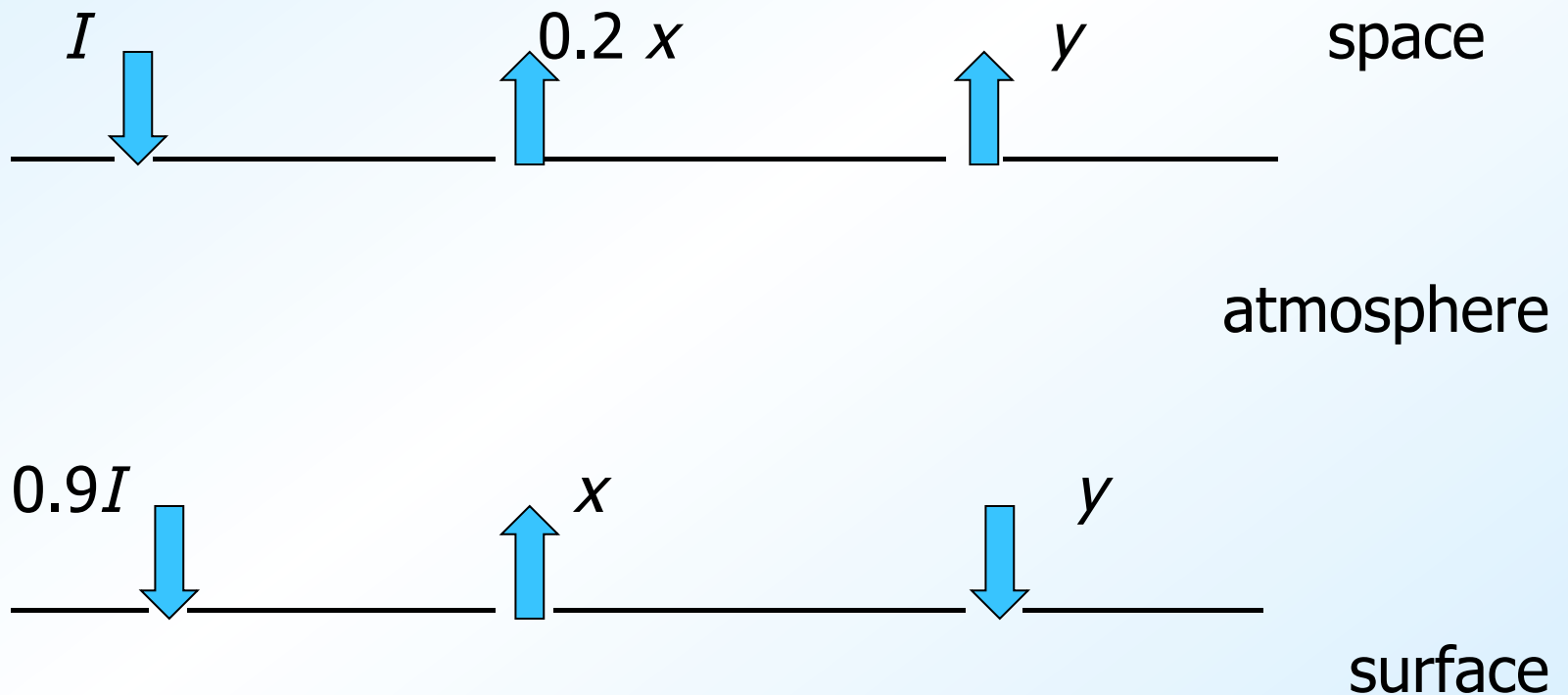
$$x = \sigma T_s^4$$

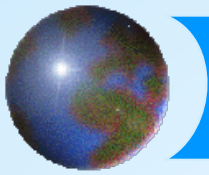
$$y = \varepsilon \sigma T_a^4 = 0.8 \sigma T_a^4$$

Why 0.8? Kirchhoff's law: $\varepsilon_\lambda = a_\lambda$



Greenhouse model





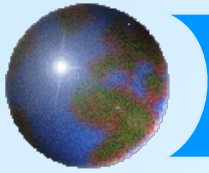
Greenhouse calculation

Balance for each level

Space: $I = 0.2x + y$

Surface: $0.9I = x - y$

Two equations, two unknowns



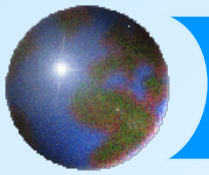
Greenhouse gas calculation

✚ Solve for x and y

$$1.9I = 1.2x$$

$$x = \frac{1.9}{1.2} I = 382.8 \text{ Wm}^{-2}$$

$$y = I - 0.2x = 165.2 \text{ Wm}^{-2}$$



Greenhouse calculation

We now have:

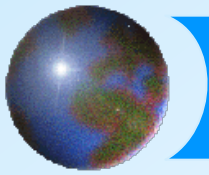
$$x = \sigma T_s^4 = 382.8 \text{ Wm}^{-2}$$

$$y = 0.8\sigma T_a^4 = 165.2 \text{ Wm}^{-2}$$

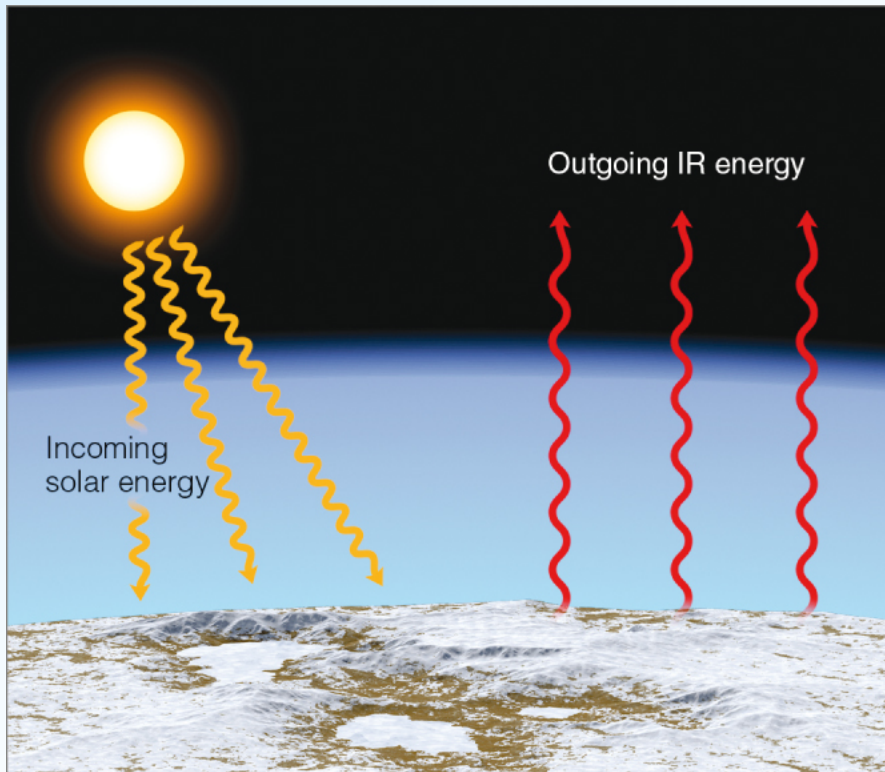
Therefore:

$$T_s = 287 \text{ K}, \quad T_a = 246 \text{ K}$$

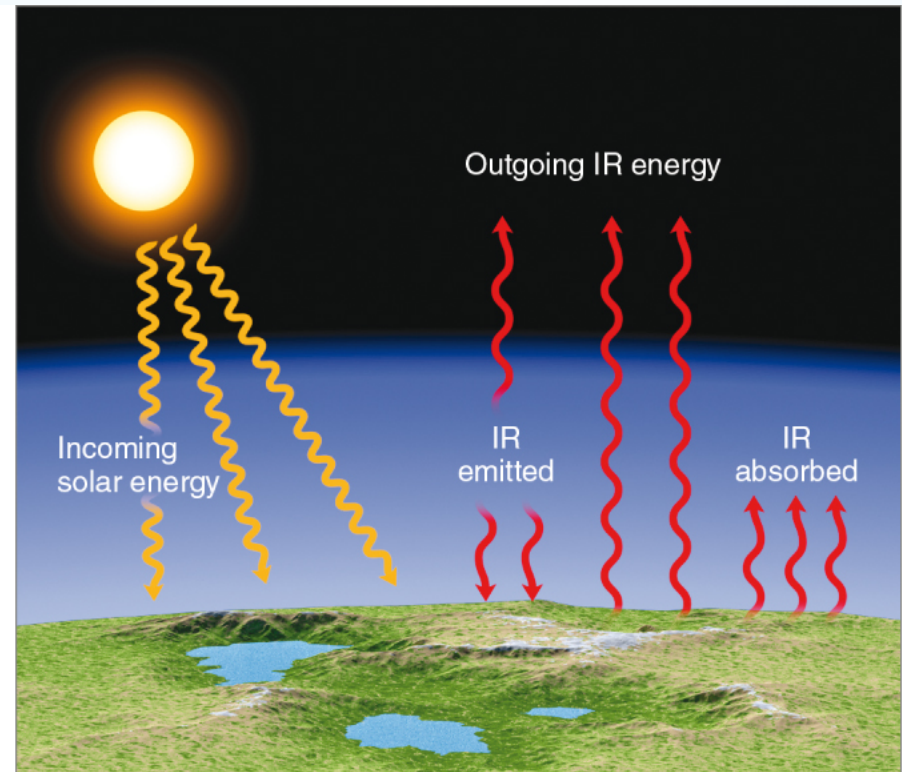
Note: Surface temperature is higher than 255 K
calculated without atmosphere



The Greenhouse Effect

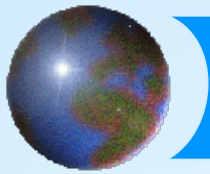


255 K



288 K

Ahrens, Fig. 2.12



Planets and atmospheres

Mars

Thin atmosphere

(Almost all CO₂ in ground)

Average temperature : - 50°C



Earth

0,03% of CO₂ in the atmosphere

Average temperature : + 15°C



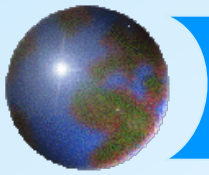
Venus

Thick atmosphere

containing 96% of CO₂

Average temperature : + 420°C





Next lecture

- ⊕ Temperature distribution
 - ⊕ Not uniform in time and place
- ⊕ Ahrens: Chapter 3