GLOBAL GENERAL CIRCULATION OF THE ATMOSPHERE

GEOG/ENST 3331 – Lecture 6 Ahrens: Chapter 10; A&B: Chapter 8; Turco: Chapter 2.4

Lecture outline

Global circulation

- One and three cell theories
- Continental effects
- Seasonality

Global circulation and surface winds

Jet streams

Scales of analysis

Scale	Size	Example
Microscale	m	Turbulent eddies
Mesoscale	km	Land/sea breeze
Synoptic scale	1 000 km	Cyclones
Global scale	5 000 km	Planetary waves

Synoptic and global scales are sometimes grouped together as 'Macroscale'

Global Circulation

Observation: There is more energy released in the polar regions than is received from the sun. The reverse is true for the equatorial region.

How can we account for this?

The atmospheric circulation must provide a poleward transport of energy

One Cell Theory

- Consider a hypothetical planet...
 - Prevailing wind from the east at the surface
- Not observed
- Need a better theory



A&B: Figure 8-2

Three Cell Theory



Three Cell Theory

Ferrel cell exists roughly between 30° and 60° in each hemisphere

- Thermally indirect, i.e. warm air sinks and cold air rises
- Surface winds travel north and producing the prevailing westerlies.



Three Cell Theory

Polar cell

- Sinking at the poles, rising at 60°.
- □ Polar easterlies are produced. Hadley







Surface zones

- Intertropical Convergence Zone (ITCZ) (Doldrums)
- Trade Winds
- Subtropical high (Horse latitudes)
- Westerlies
- Polar Front
- Polar easterlies



The ITCZ is observable as a band of clouds extending from northern South America into the Pacific



Axial tilt in Earth's orbit



Seasons and continents

- Impact of seasons
 - General circulation features shift north in JJA, south in DJF
 - This is most evident in the Northern Hemisphere

- Impact of continents
 - Land masses warm up and cool down faster as the seasons change

Global Precipitation Pattern Produced by the General Circulation







Upper level flow

- □ Much less friction; winds are geostrophic
 - There is much less meridional heat transport
 - Strong zonal heat transport

- Impacts of seasons and continents
 - Circulation still shifts with the seasons
 - Land/sea contrast less evident





Lecture outline

- Global circulation
- Global circulation and surface winds
- Jet streams
 - Polar jets
 - Rossby waves

Jet Streams

Swift flowing current of air

- Thousands of km long, a few hundred km wide, a few km thick and 10-15 km above the surface.
- Speed ranges from 150 to 300 km/h.
- Jets occur at the divisions of the three cells.



Polar Jet



- For midlatitude regions the polar jet is more important.
- Boundary between cold and warm air.
- Surface features, such as air masses and storms, tend to follow the direction of the upper level jet stream.

Ahrens: Active Fig. 10.10

Polar Jet

- Strong temperature gradient between Polar and Ferrel cells leads to strong horizontal pressure gradient.
- Geostrophic balance causes a strong wind parallel to isobars.



Ahrens: Fig. 10.12

Rossby Waves

Upper air flow

At any given time there are 3-6 planetary waves or longwaves.



Ahrens: Fig. 12.8

Planetary Waves

- □ Slow moving
 - Can be stationary for months
 - Migrate slowly west to east
 - Sometimes east to west
- □ Winter
 - Fewer, longer, stronger





Shortwaves

- Rossby waves that travel eastward along the longwaves.
- □ These are smaller, shorter, faster-moving disturbances.
- Shortwaves become stronger near the troughs and weaker near the ridges of longwaves.

Shortwaves and longwaves



Ahrens: Active Fig. 12.9

Rossby Waves

Surface flow follows upper atmosphere flow

Mild in Yukon

Cold over plains, central US

September 22, 1995

A&B: Figure 8-13(a)



Coming up

- Atmospheric observation
- Weather prediction
- Numerical modelling