



Global Circulations

GEOG/ENST 2331 – Lecture 15 Ahrens: Chapter 10



Last lecture

Microscale (turbulence)
Mesoscale (land/sea breeze)
Synoptic scale (monsoon)
Global scale (3 cell circulation)









Upper level flow

- Much less friction; winds are geostrophic
 There is much less *meridional* heat transport
 Strong *zonal* heat transport
- Impacts of seasons and continents
 Circulation still shifts with the seasons
 Land/sea contrast less evident





Ahrens: Figure 10-9





Jet Streams

Swift flowing current of air

- Thousands of km long, a few hundred km wide, a few km thick and 10-15 km above the surface.
- Speed ranges from 150 to 300 km/h.
- Jets occur at the divisions of the three cells.





Polar Jet



- For midlatitude regions the polar jet is more important.
- Boundary between cold and warm air.
- Surface features, such as air masses and storms, tend to follow the direction of the upper level jet stream.

Ahrens: Active Fig. 10.10



Polar Jet

- Strong temperature gradient between Polar and Ferrel cells leads to strong horizontal pressure gradient.
- Geostrophic balance causes a strong wind parallel to isobars.



Ahrens: Fig. 10.12

The "Dishpan Experiment"









Rossby Waves

 Also known as longwaves or planetary waves

Upper air flow

 At any given time there are 3-6 Rossby waves.

C.G. Rossby – famous meteorologist in early 20th Century.

Ahrens: Fig. 12.8



Rossby Waves

- Slow moving
 - Can be stationary for months
 - Migrate slowly west to east
 - Sometimes east to west

Winter

 Waves are fewer, longer, stronger





A&B: Figure 8-11







Shortwaves

- Small disturbances or ripples embedded in Rossby waves,
- Faster-moving and travel eastward along the Rossby waves.
- Shortwaves become stronger near the troughs and weaker near the ridges of Rossby waves.



Shortwaves and Longwaves (Rossby waves)



Ahrens: Active Fig. 12.9

Global Ocean Circulation



Ekman Spiral

Wind drag pushes the surface waters 45° to the right (in the NH)

Deeper water is dragged along at successively greater angles

Around 100 m direction is opposite surface winds and current dies out

Surface water Wind 45° Wind 100 m Depth Average transport

Ahrens: Fig. 10.17

Upwelling

Offshore winds, and even parallel-to-shore winds drag away the surface water and cause upwelling of **cold**, nutrient-rich water from below.

A&B: Figure 8-17





ENSO

El Niño / Southern Oscillation

- 4 to 7 year oscillation in tropical Pacific Ocean temperatures and associated changes in atmospheric circulation
- Multiple causes: linked to changes in the Asian monsoon season, long-range transport of subsurface waves, etc.
- Irregular cycle
- Once in process weather predictions for the next season or two can be made, often with considerable accuracy.



The "Normal" Walker Circulation



A&B: Figure 8-27

Walker Circulation

- Easterly trade winds
- Low pressure, rising air and heavy rain over western Pacific
- Winds push surface water away from South America, causing cold water to well up from below
- Particularly intense years are called La Niña



Ahrens: Fig. 10.19

El Niño

- Winds weaken or reverse
- High pressure, subsiding air, drought over western Pacific
- Upwelling in eastern Pacific stops, surface temperatures rise









Average sea surface temperature differences from normal during the November through March period for eight El Niño episodes. A&B: Figure 8-29

Teleconnections

Relationships between climatic patterns at widely separated locations

ENSO affects upper level flow and position of jet streams.

Polar jet stream Warm Cool Wet Wet Jet stream (a) El Niño winter conditions Polar jet Cold Н Wet Wet Warm Southern branch of polar jet Dry Warm (b) La Niña winter conditions

Ahrens: Fig. 10.23 Impacts of El Niño

A&B: Figure 8-35







Blocking (associated with ENSO)

Example of an omega block over western North America – Warmer and drier within the block

05/15/2006 1Burte 05/15/2006 DOUTC 018HR TO HON

060515/1800V018 HAM 500 HB HGT, GEO ABS VORTICITY

Figure May 2006



ENSO Index

Strong positives are El Niño Strong negatives are La Niña





North Atlantic Oscillation (NAO)



Positive

Ahrens: Fig. 10.27

Negative

Arctic Oscillation (AO)





Ahrens: Fig. 10.28



Next lecture

Air Masses and FrontsAhrens: Chapter 11